

**Lauder, New Zealand within Network for
Detection of Atmospheric Composition
Change (NDACC)**

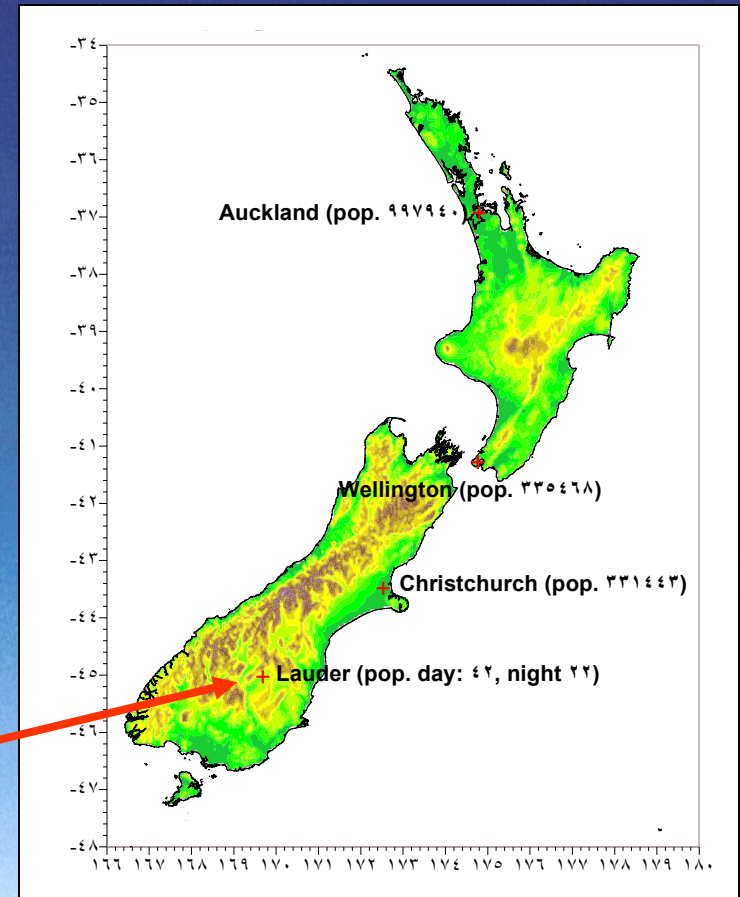
Presented at

**Meeting on Implementation of GCOS
Reference Upper Air Network
(GRUAN)**

Lindenberg, Germany, 26-28 February 2008

NIWA's Lauder Site

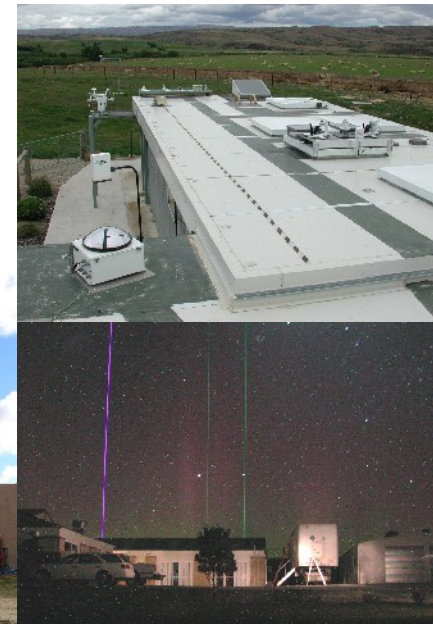
**45.0°S, 169.7°E
370m a.s.l.**



“Lauder Within the NDACC” – we will look at:

- 1. Location – Geography and Climate**
- 2. NDACC Measurements**
- 3. Links with GRUAN goals:**

- Measurements**
- Climate Modelling**
- Research Outputs**

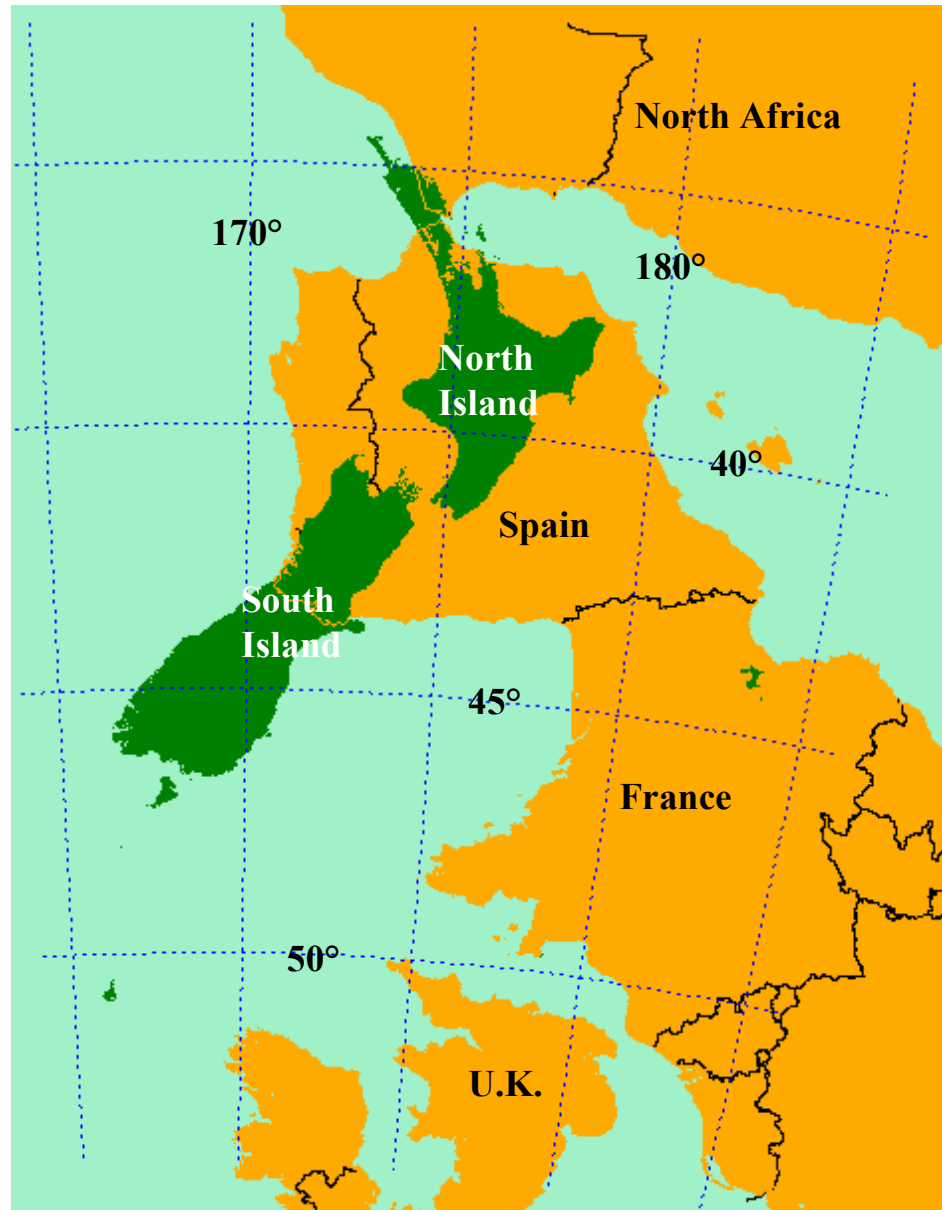


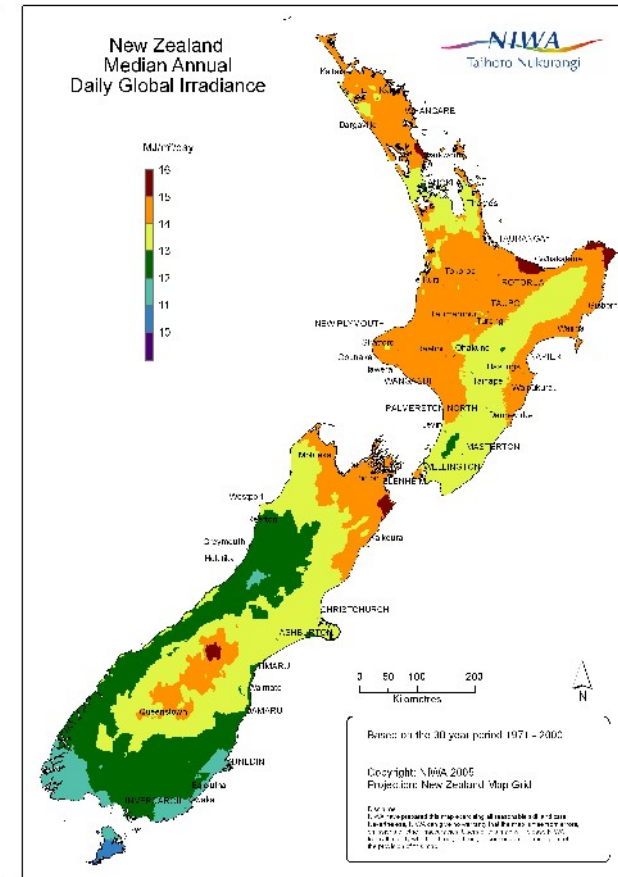
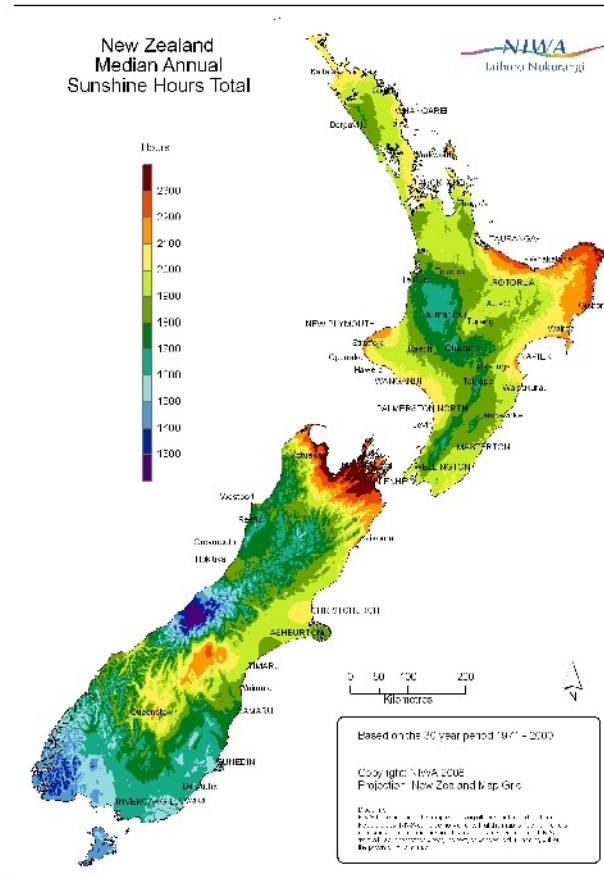
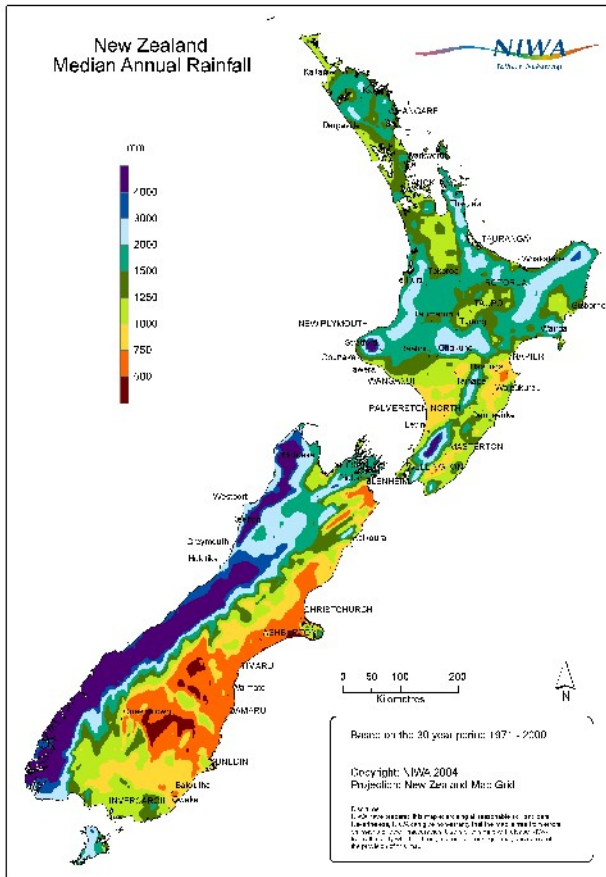
New Zealand, a small country close to Australia and Antarctica



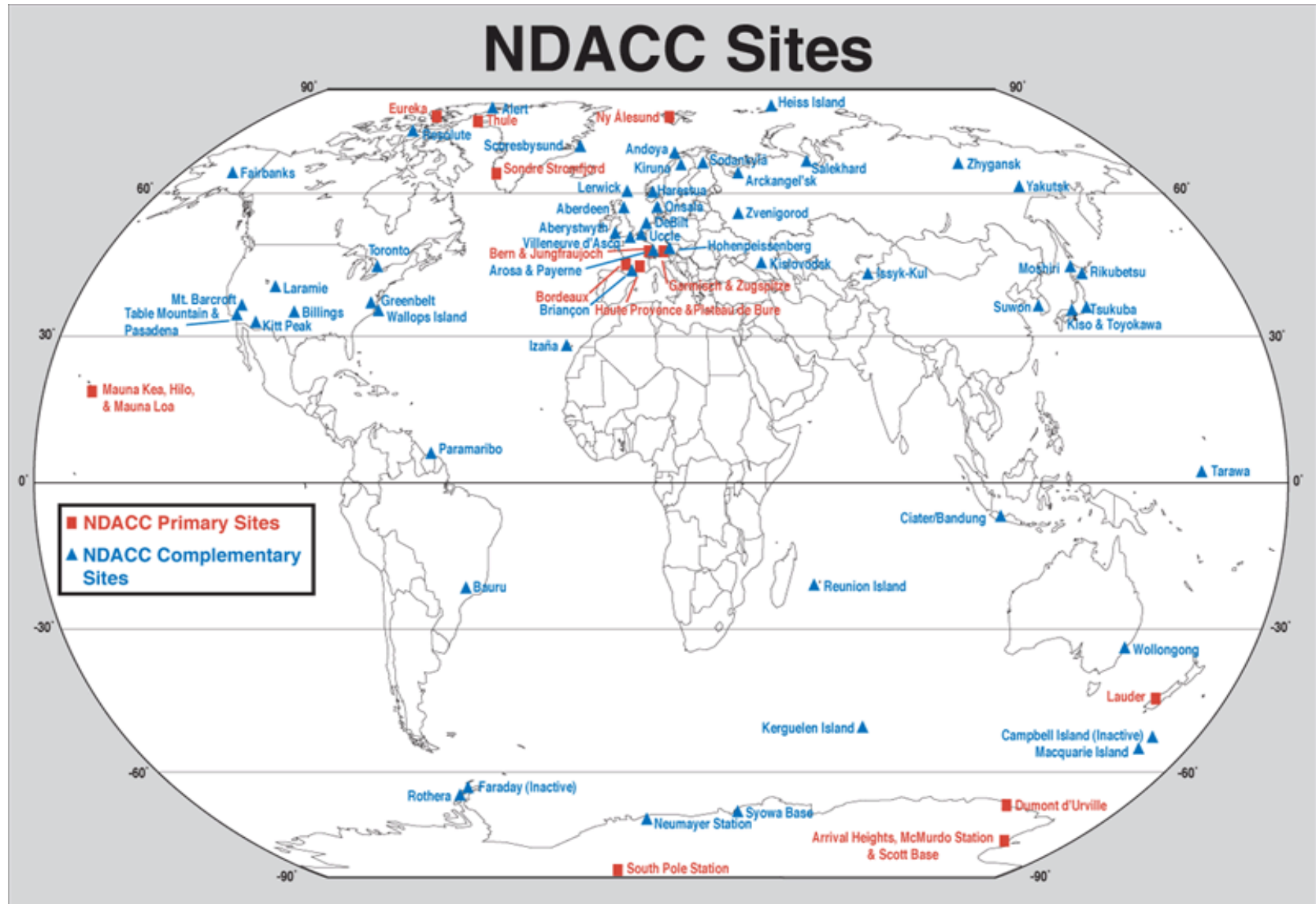
McKenzie et al., JGR 103, D22, pp28,785-28,792, 1998.
Image from J Descloitres, MODIS Rapid Response Team
at NASA GSFC, as published in NIWAClimate Update
No. 50

New Zealand and its Antipodes

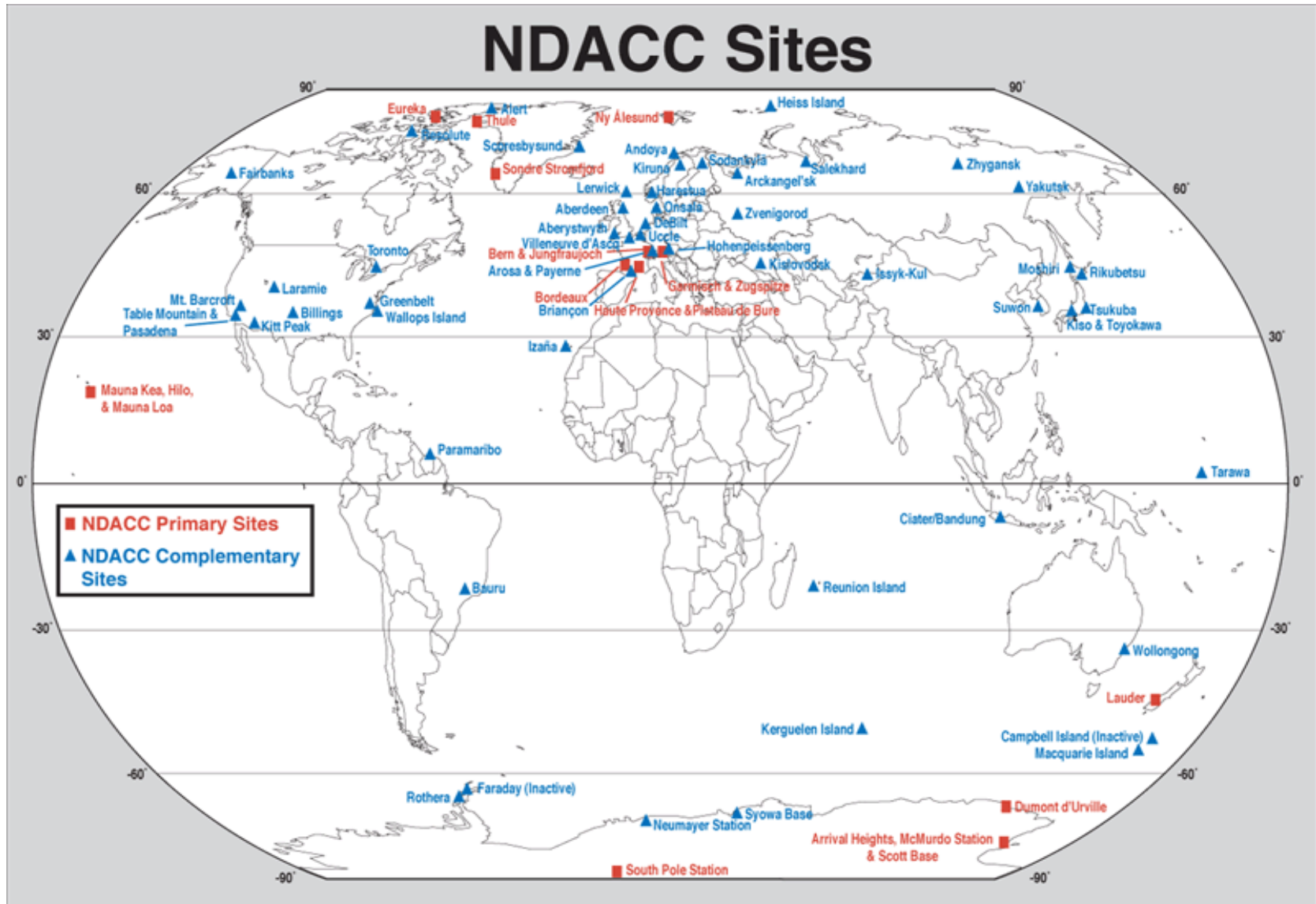




The international Network for the Detection of Stratospheric Change (NDSC) was formed to provide a consistent, standardised set of long-term measurements of atmospheric trace gases, particles, and physical parameters via a suite of globally distributed sites.



The international Network for the Detection of Atmospheric Composition Change was established to provide a consistent, standardised set of long term measurements of atmospheric trace gases, particles, and physical parameters via a suite of globally distributed sites.



Reference Upper Air Network (GRUAN), Lindenberg, Germany, 26-28 February 2008: Lauder, New Zealand within NDACC

The primary NDACC instruments and measurements are:

- **Ozone Sondes and Dobson/Brewer** (Profiles to $> 30\text{km}$ & total column ozone)
- **Ozone lidar** (vertical profiles of ozone from the tropopause to at least 40 km altitude; in some cases tropospheric ozone will also be measured)
- **Temperature lidar** (vertical profiles of temperature from about 30 to 80 km)
- **Aerosol lidar** (vertical profiles of aerosol optical depth in the lower stratosphere)
- **Water vapor lidar** (vertical profiles of water vapor in the lower stratosphere)
- **Ozone microwave** (vertical profiles of stratospheric ozone from 20 to 70 km)
- **H₂O microwave** (vertical profiles water vapor from about 20 to 80 km)
- **ClO microwave** (vertical profiles of ClO from about 25 to 45 km , dep. on latitude)
- **Ultraviolet/Visible spectrograph** (column abundance of ozone, NO_2 , and, at some latitudes, OClO and BrO)
- **Spectral UV Radiometer** (irradiances to track effects of Atmospheric Change)
- **Fourier Transform Infrared spectrometer** (column abundances of a broad range of species including ozone, HCl , NO , NO_2 , ClONO_2 , and HNO_3)

Measurements at the Lauder NDACC Station, Dec 2007

Instrument & Period	Parameter	Cooperating Institutions	Comments
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Ozone Dobson/Sonde, Lauder

Dobson Ozone 1987 Jan -	Column ozone	NOAA/GMD, Boulder, USA	
Balloon sonde (ECC) 1986 Aug -	Ozone, T/P, humidity and wind profiles	NOAA/GMD, Boulder, USA	Weekly throughout the year. 0 – 32 km

FTIR Trace Gas Spectroscopy, Lauder

Mid IR Interferometer (Bruker) 1990 -	Column HCl, HNO ₃ , ClONO ₂ , HF, CFCs, CO & GHGs		Mid-IR (windows in 2-12μ region). SFIT2 profiles
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Total Column Carbon Observing Network, Lauder

Near IR Interferometer (Bruker) 2003 -	Column CO ₂ , CO, CH ₄ , N ₂ O	NASA, JPL, UC, Univ. Wollongong	Total Column Carbon Netw. OCO Validation
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UV-Vis Trace Gas Spectroscopy, Lauder

UV-Vis DOAS spectrometers 1980 Dec -	Column NO ₂ Column BrO Zen (1994) + Sun (2001) + Off-axis (04)	NOAA, Univ. Nagoya, Swedish IRF Aus. Ant. Div.	6 Lauder systems 4 O/Seas systems
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Instrument & Period	Parameter	Cooperating Institutions	Comments
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Ozone/Aerosol Lidars, Lauder

Aerosol & temperature lidar. 1993 Dec -	Aerosols 5-30 km Temp. 30-70 km	IROE, Florence, Italy University of Lyon, France	Nd-YAG 532 nm, 355 nm
Aerosol lidar. 1992 Nov -	Aerosol profile 3-30 km	Met. Research Instit., Tsukuba, Japan	Nd-YAG 532 nm
Ozone lidar 1994 Dec -	Ozone profiles 8-45 km	RIVM, Bilthoven, Netherlands	Excimer laser plus Raman cell

Microwave Radiometers, Lauder

microwave radiometer 1993 -	Ozone profiles 20-65 km	University of Massachusetts, USA	110 GHz and 6 hour integration
Microwave radiometer 1994 -	H ₂ O 40-80 km profiles	NRL, Washington DC, USA	22 GHz and 1 week integrat. times

Spectral UV, Lauder

UV Spectrometer 1980 & 1988 obs. 1989 Nov -	Global and scattered UV Spectral irradiance	NOAA/GMD, Boulder, USA. BOM, Australia	285-450 nm 3 Lauder systems 5 O/Seas systems
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NDACC Measurement Intercomparisons

A fundamental requirement of NDACC operation is involvement in regular intercomparisons of instruments and processing algorithms. These ensure measurement consistency across the global network as well as providing close interaction of experimenters from all organisations making NDACC measurements. Considerable improvements in accuracy and stability have been gained as a result.

Recent Intercomparisons by Working Group, are:

Dobson: Lauder, 2001

Lidar: Ozone Lidar, Lauder, 2002

Ozone Sondes: Balloon Experiment on Standards for Ozone Sondes (BESOS), 2004

FTIR: Kiruna, ?????

UV/Visible: Andoya, Norway, 2003

Spectral UV: Table Mountain, Boulder, 2003

Ozone and Hygrometer Sonde measurements + Dobson Ozone

Weekly Sonde measurements are made using two different systems on alternate weeks:

- RS92-SGP Radiosonde with ECC-1Z Ozone Sensor (1986 on). NOAA copy.
- RS80-15H Radiosonde with ECC-1Z Ozone Sensor and NOAA Frost Point Hygrometer.

Both use a GPS to record winds.

Total Column ozone is measured up to 5 times/day in clear sun conditions. The Lauder Dobson Spectrophotometer of the part of International Dobson Network run by NOAA.



Lauder Sonde with RS92-SGP and NOAA Hygrometer Specifications

(GRUAN Requirements in Red)

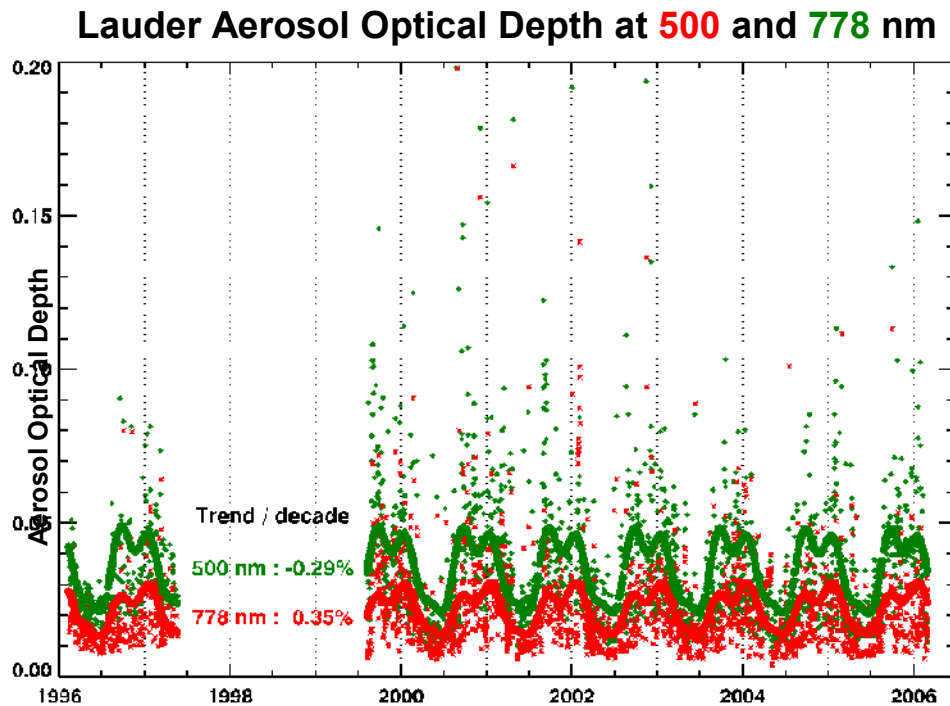
Pressure	Type:	Barocap silicon sensor
	Range:	1080 to 3 hPa
	Resolution:	0.1 hPa 0.01 hPa
	Total Uncertainty in Sounding ¹	1.5 hPa (1080 – 100 hPa) 0.1 hPa 0.6 hPa (100 – 3 hPa)
	Repeatability ²	0.4 hPa (1080 – 100 hPa) 0.3 hPa (100 – 3 hPa)
	Reproducibility ³	0.5 hPa (1080 – 100 hPa) 0.3 hPa (100 – 3 hPa)
Temperature	Type:	F-Thermocap capacitive wire
	Range:	+60°C to -90°C
	Resolution:	0.1°C
	Accuracy:	
	Total Uncertainty in Sounding ¹	0.5°C Accuracy: 0.1°K Tropo 0.2°K Strato
	Repeatability ²	0.15°C Long term stab: 0.05°K
	Reproducibility ³	0.2°C (1080 – 100 hPa) 0.3°C (100 – 20 hPa) 0.5°C (20 – 3 hPa)
	Lag:	< 0.4 s (1000 hPa) < 1 s (100 hPa) < 2.5 s (10 hPa)

Wind	Type:	Code-correlating GPS receiver, 12 channels
	Accuracy:	
	- Horizontal position	10m
	- Vertical position	20m Vert Res. 50 m tropo 250 m strato
	- Wind velocity	0.2m/s 0.5 m/s
Water Vapour	Type:	Chilled Mirror Hygrometer (NOAA)
	Accuracy	Approx 10% 2%
Ozone Sonde	Type:	Coulometric, employing a electrochemical concentration cell with platinum electrodes, and a Teflon gas sampling pump
	Estimated accuracy:	< 10% 5%
	Noise:	< 1% of full scale
	Sensitivity:	2 to 3 parts per billion by volume

Radiation and Aerosol Lidar

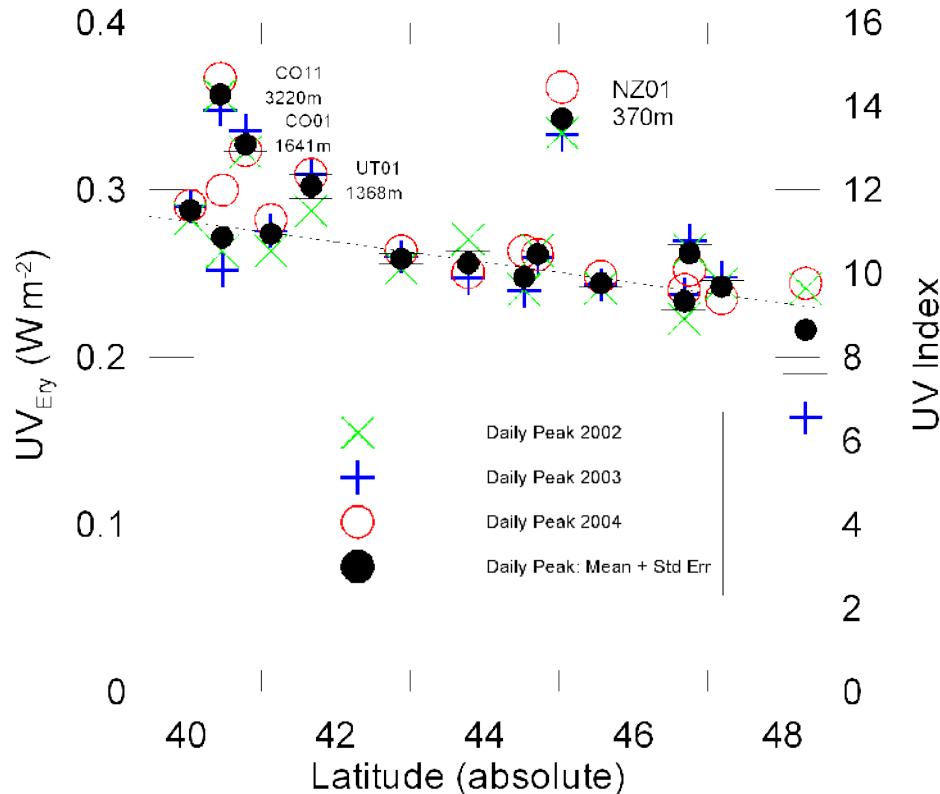
Lauder is a BSRN site in collaboration with Bureau of Meteorology, Melbourne.

- BSRN measures the downward components of solar and Thermal irradiance from a globally distributed surface based network of 35 sites.
- Ancillary BSRN data at Lauder is: meteorological observations, aerosol optical depth and Spectral UV.
- Clean atmosphere at Lauder shown in AOD record:



Spectral UV Measurements

Peak UV at 45 S (Lauder, NZ)
compared with sites in North
America of comparable latitude.



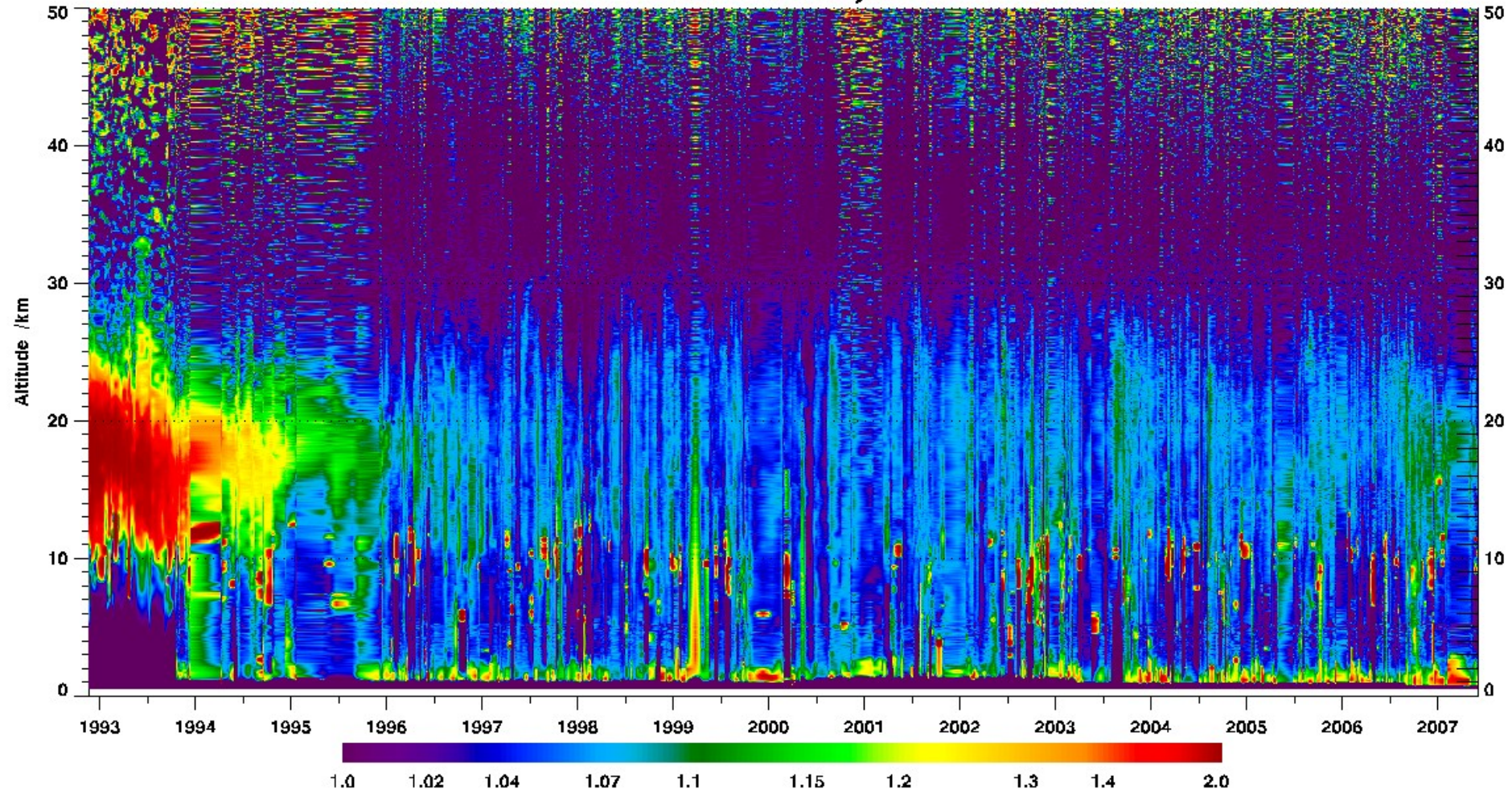
NIWA UV Radiometer measuring
Spectral Irradiance, 285 – 450 nm.
Accuracy $\pm 3\%$, traceable to NIST.

Aerosol Lidar Measurements

The results shown are from the MRI Japanese Lidar processed using algorithms developed at Lauder.

The Lauder analysed backscatter ratio is shown, interpolated between measurements.

MRI Lidar Reanalysis



Synopsis

- Future climate predictions uncertain, in part because the carbon budget cannot be balanced.
- A new generation of satellite measurements (OCO, GOSAT) will be coupled with high precision, ground-based column density measurements, to determine the surface to atmosphere flux of CO₂ and CH₄.
- Carbon flux presently inferred from a network of surface measurements, interpreted within atmospheric transport models, but surface measurements alone have limited success in inferring carbon flux to the atmosphere.
- Column measurements are much less sensitive to uncertainties in vertical mixing, but respond similarly to surface flux. Thus the combination of column and surface measurements is a powerful constraint on model transport, and enables confident determination of flux.
- Satellite column measurements complement ground-based observations by providing dense global coverage.
- Measurement and modeling capabilities have now advanced to the point that it is possible to exploit this synergy between surface, column, and satellite measurements.

- Hi-resolution FTS in near-IR solar absorption making precision column measurements of CO₂, CH₄, CO & N₂O.

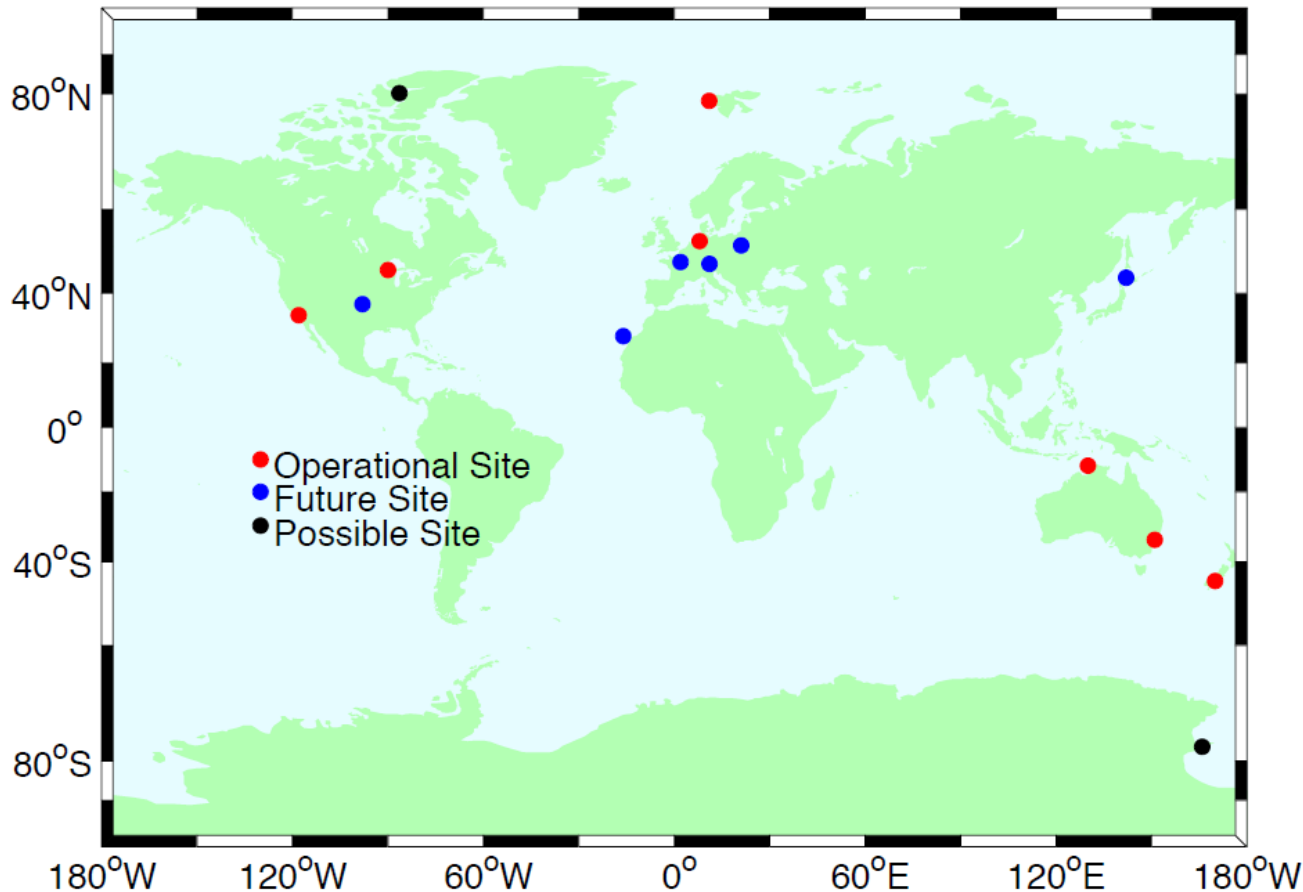
- Continuous in-situ measurements of CO₂ at all sites.

- Continuous in-situ measurements of CH₄, CO, N₂O at selected sites (Lauder, Darwin, others TBD).

- An in-situ precision CO₂ analyser (Licor 7000) will be installed at Lauder soon.



Total Carbon Column Observing Network (TCCON) Sites



Operational

Bremen, Germany

Ny-Alesund, Norway

Wollongong, Australia

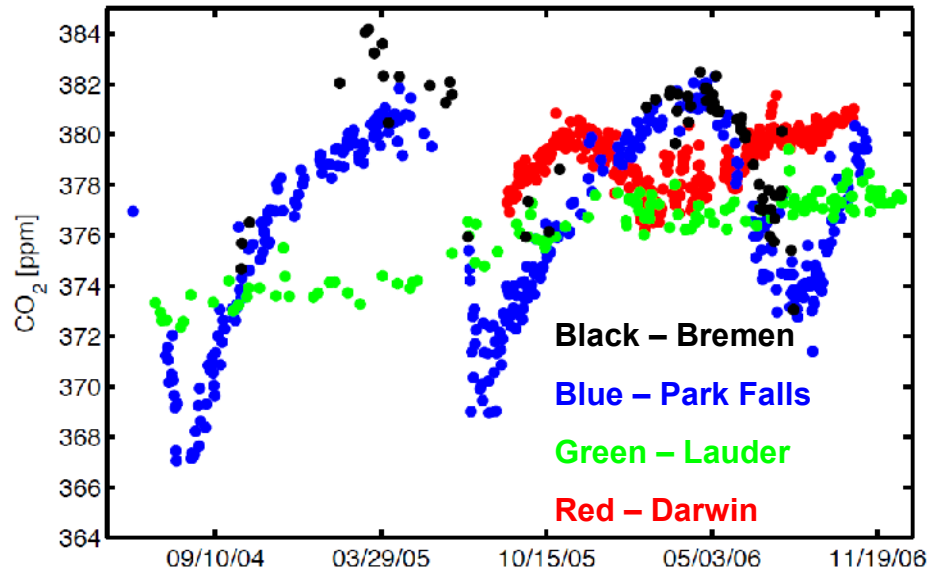
Darwin, Australia

Lauder, New Zealand

**Park Falls, Wisconsin,
USA**

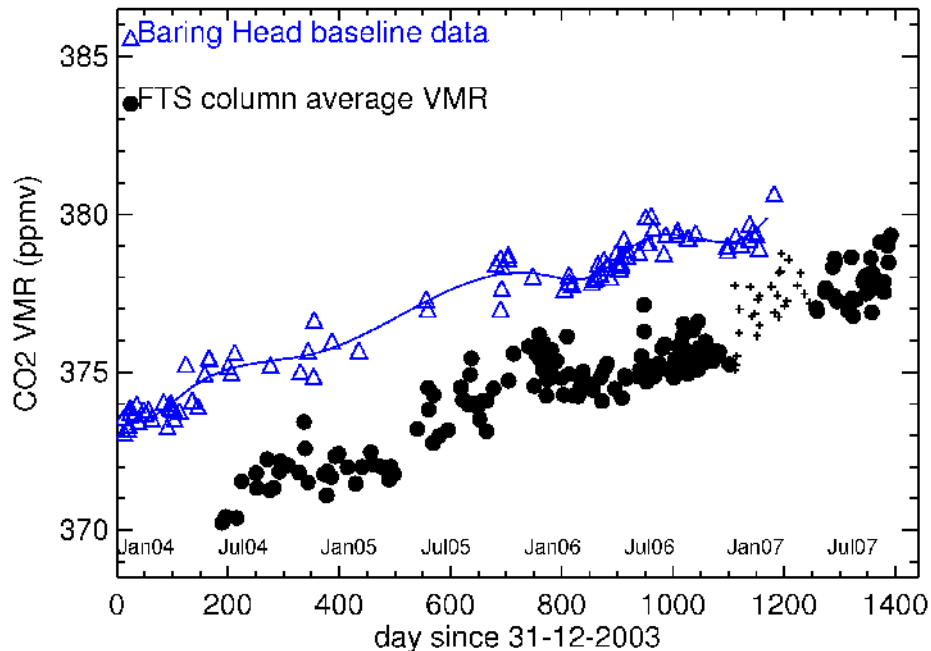
CIT, Pasadena, USA

Examples of TCCON data



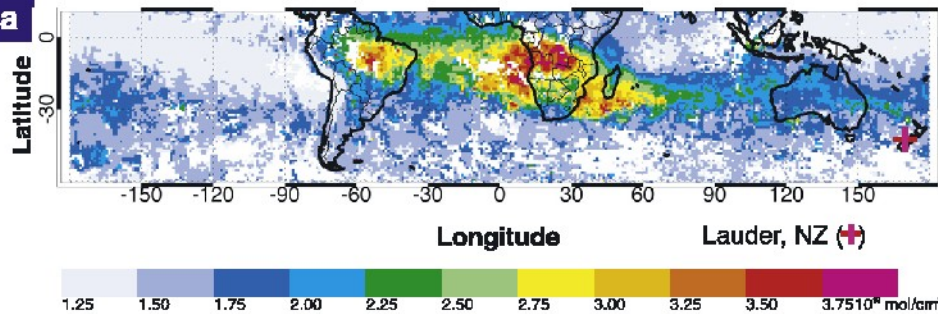
Average Vol. Mix. Ratios are derived from the greenhouse gas column amount and the O₂ column amount, obtained from the same measurement spectra.

The daily average CO₂ precision is ≤ 0.6 ppm.

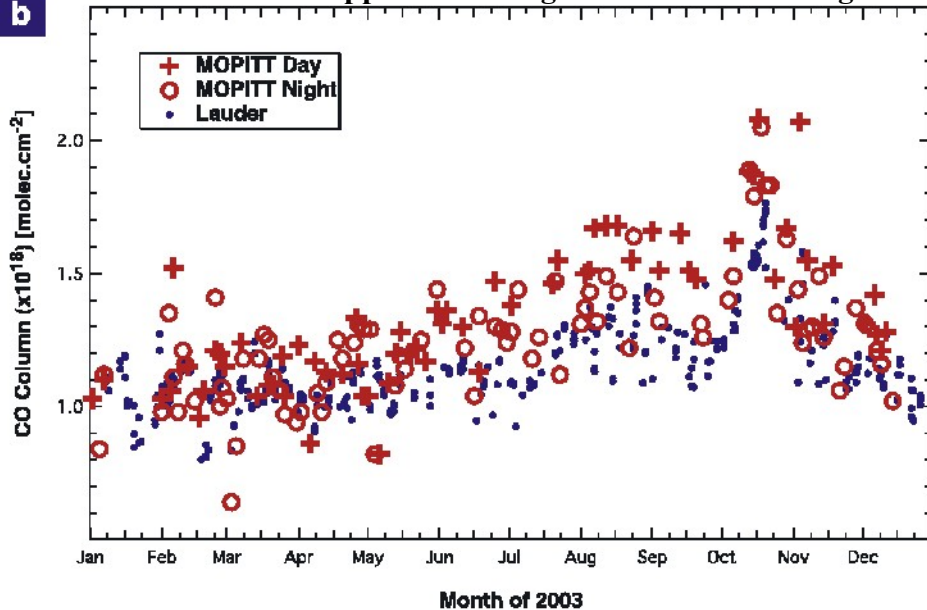


Examples of the use of Mid-IR FT-Spectroscopy to measure atmospheric composition at Lauder and Validate Satellite measurements

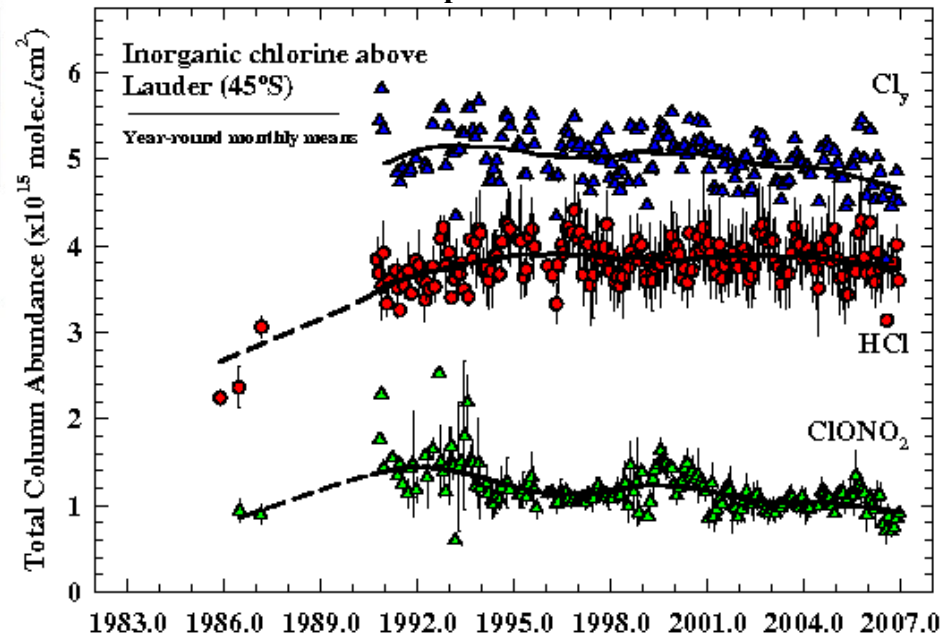
a Mean distribution of MOPITT CO total column 1-9 October 2003



b Ground Based + Moppit CO during 2003 at Lauder using FTIR

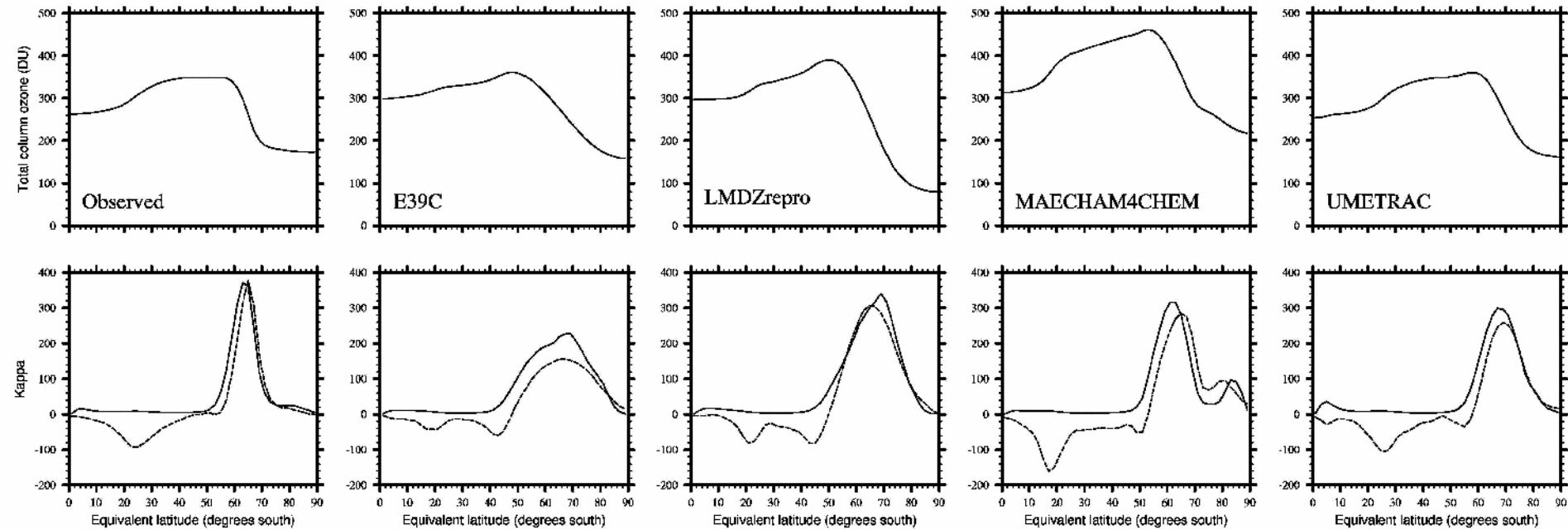


Time Series of Stratospheric Chlorine above Lauder



Coupled Chemistry Climate Modelling at Lauder

Comparison of equivalent latitude zonal mean ozone total column and meridional impermeability from observations and four coupled chemistry-climate models.



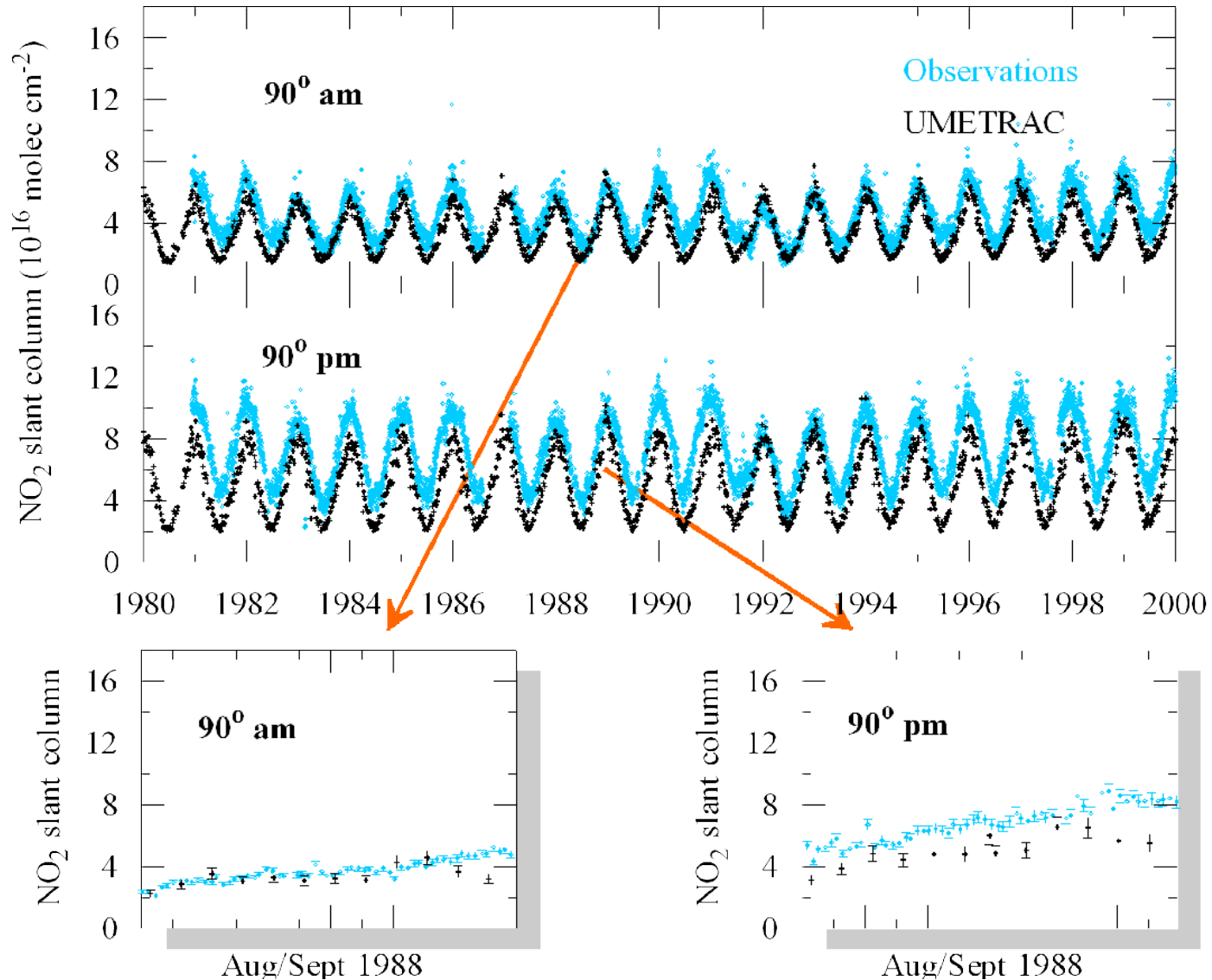
Top plots shows the sharp gradients at the edge of the ozone hole.
Solid line in the bottom plots = kappa (meridional impermeability).
[Kappa is a diagnostic of the strength of the dynamical mixing barrier.]

Dashed lines = the negative of the total column ozone gradient.
The bottom plots demonstrate that:

- (b) there is a narrow mixing barrier at the edge of the polar vortex. All the models produce kappa peaks that are too wide and generally a little too weak.
- (b) Gradients in total column ozone are strongly correlated with kappa (both in the real atmosphere and in models).
- (c) The models representation of the ozone gradients at the edge of the polar vortex are too weak and this is strongly related to the modelling of dynamics at the vortex edge.

EMETRAC:
Unified Model with
Eulerian Transport
and Chemistry

Measured NO₂ v CCM (Unified Model with Eulerian Transport and Chemistry)



Struthers, et. al., Atmos. Chem. Phys., 4, 2227–2239, 2004

Lauder Staff

Scientists: 8

Engineers/Technicians: 6

Administration/Computing: 2

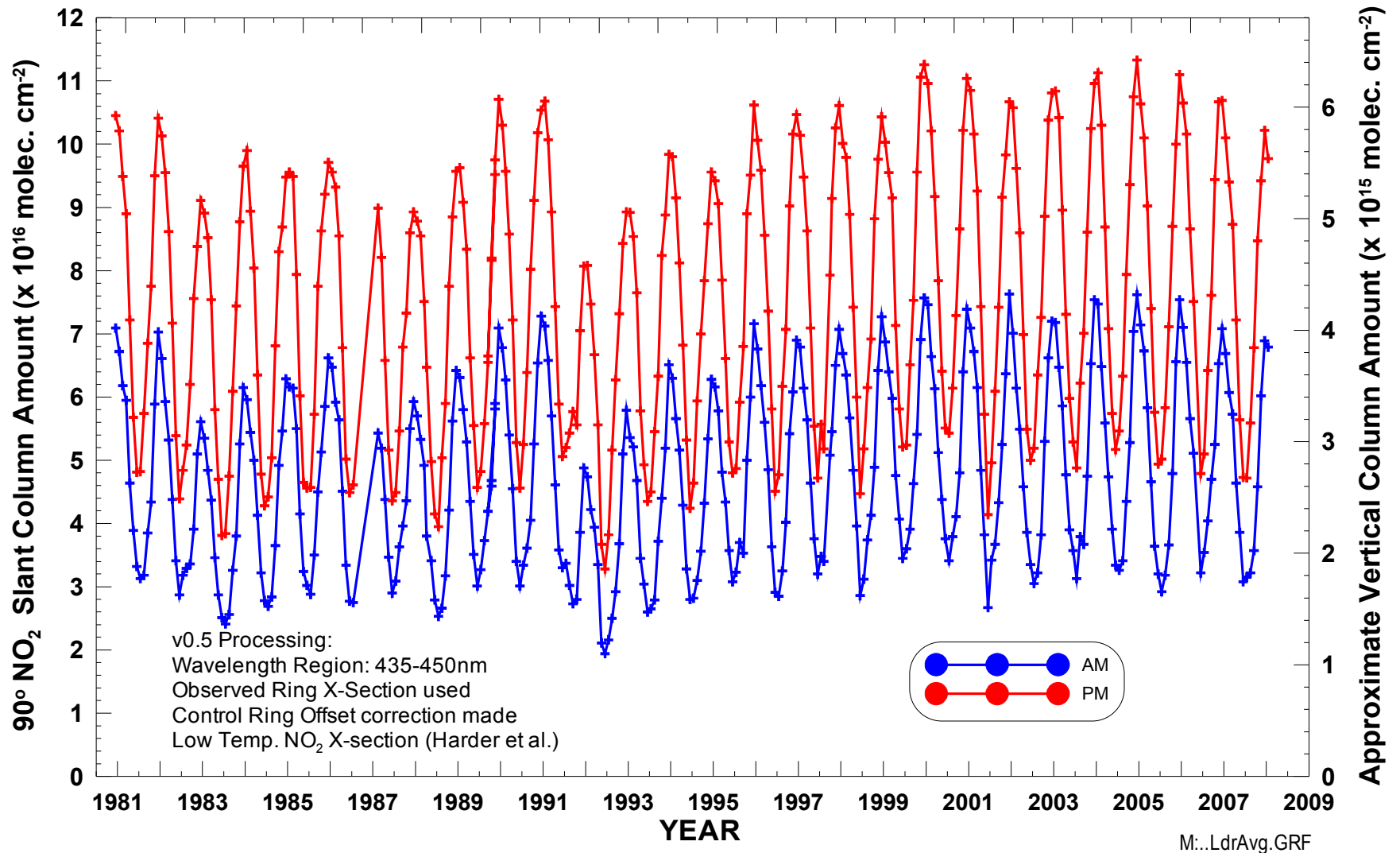
Students: typically 2 PhD students + 2 overseas students gaining practical experience.

Published Research Papers 2003-2007

Lauder First Author: 26

Lauder Co-Author: 91

LONG TERM NO₂ MEASUREMENTS OVER LAUDER, NZ (45° S, 170° E)



Past and future simulations of NO₂ from a coupled chemistry-climate model in comparison with observations

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Received: 8 June 2004 – Published in Atmos. Chem. Phys. Discuss.: 20 August 2004

Revised: 15 November 2004 – Accepted: 15 November 2004 – Published: 22 November 2004

Abstract. Trends in NO₂ derived from a 45 year integration of a chemistry-climate model (CCM) run have been compared with ground-based NO₂ measurements at Lauder (45° S) and Arival Heights (78° S). Observed trends in NO₂ at both sites exceed the modelled trends in N₂O, the primary source gas for stratospheric NO₂. This suggests that the processes driving the NO₂ trend are not solely dictated by changes in N₂O but are coupled to global atmospheric change, either chemically or dynamically or both. If CCMs are to accurately estimate future changes in ozone, it is important that they comprehensively include all processes affecting NO_x (NO+NO₂) because NO_x concentrations are an important factor affecting ozone concentrations. Comparison of measured and modelled NO₂ trends is a sensitive test of the degree to which these processes are incorporated in the CCM used here. At Lauder the 1980–2000 CCM NO₂ trends (4.2% per decade at sunrise, 3.8% per decade at sunset) are lower than the observed trends (6.5% per decade at sunrise, 6.0% per decade at sunset) but not significantly different at the 2σ level. Large variability in both the model and measurement data from Arival Heights makes trend analysis of the data difficult. CCM predictions (2001–2019) of NO₂ at Lauder and Arival Heights show significant reductions in the rate of increase of NO₂ compared with the previous 20 years (1980–2000). The model results indicate that the partitioning of oxides of nitrogen changes with time and is influenced by both chemical forcing and circulation changes.

1 Introduction

It has been recognised for some time that the reactive species NO_x are important in the altitude range from approximately 20 km to 35 km in determining the concentration of stratospheric ozone (Crutzen, 1970). NO_x destroys ozone through the catalytic cycle shown in Reactions (1) and (2).



In the lower stratosphere (approximately 10 km to 20 km) the most significant influence of NO_x is its interaction with the ClO_x and BrO_x ozone loss cycles via the formation of reservoir species ClONO₂ and BrONO₂. NO₂ also reacts with OH (Eq. 4), reducing the HO_x concentration and thus inhibiting the HO_x catalysed destruction of ozone.

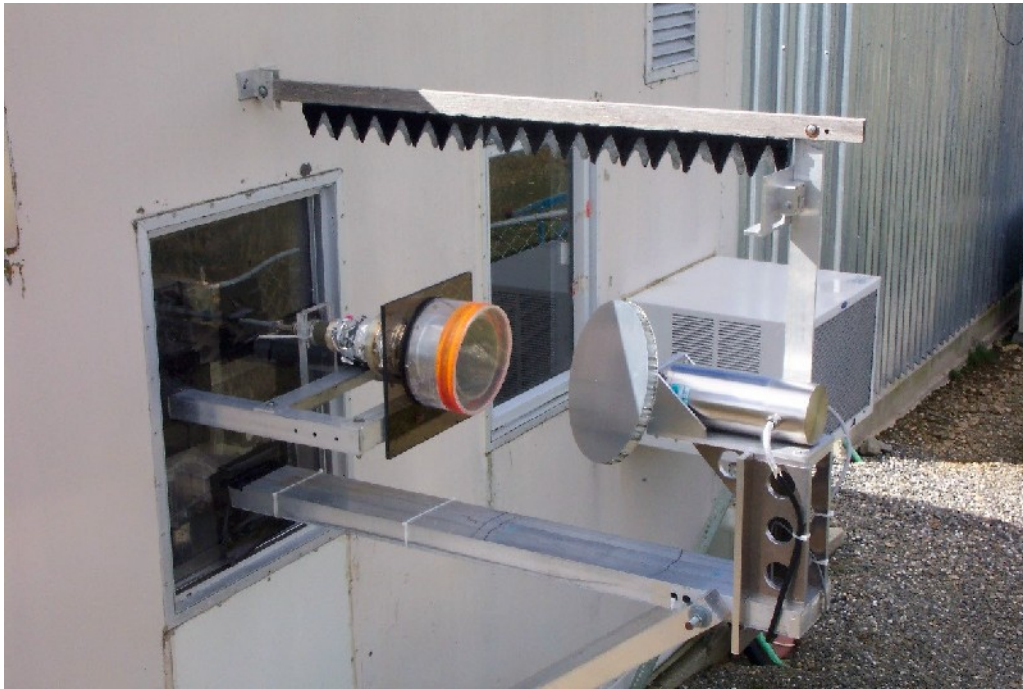


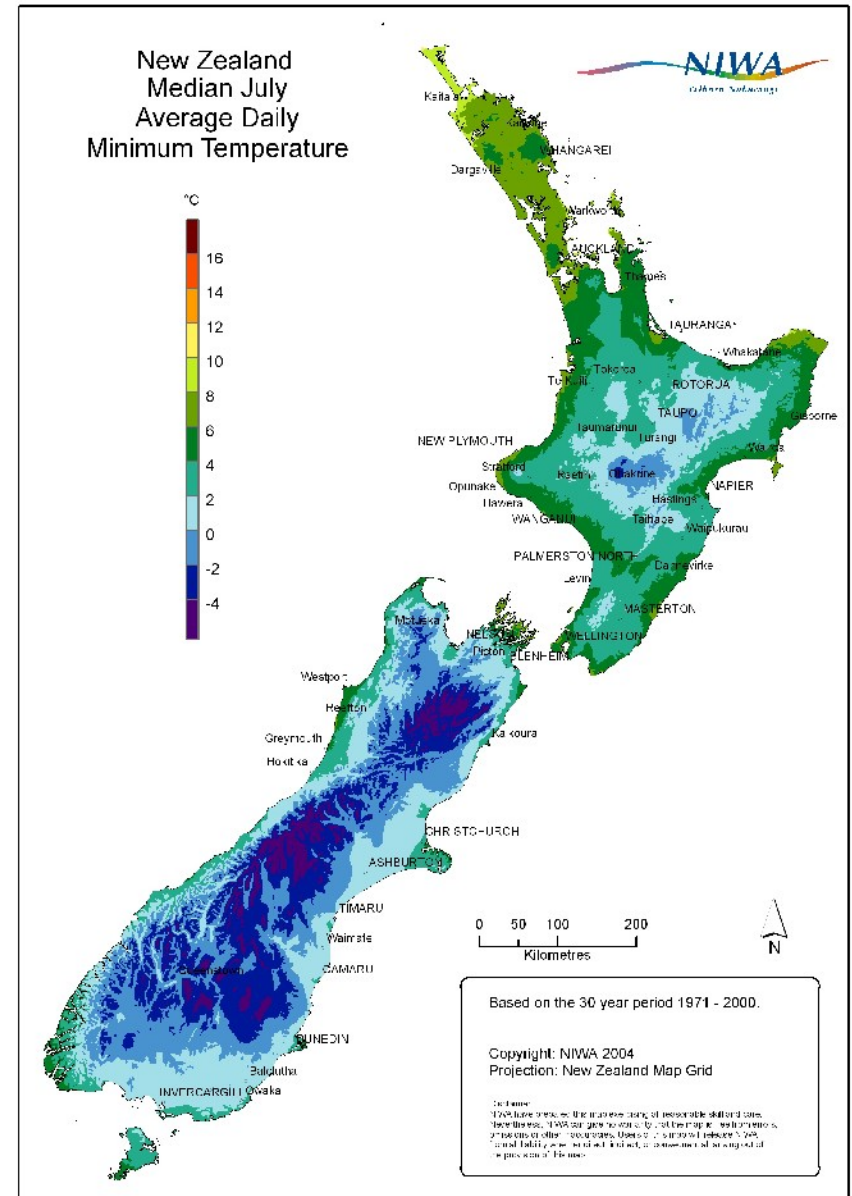
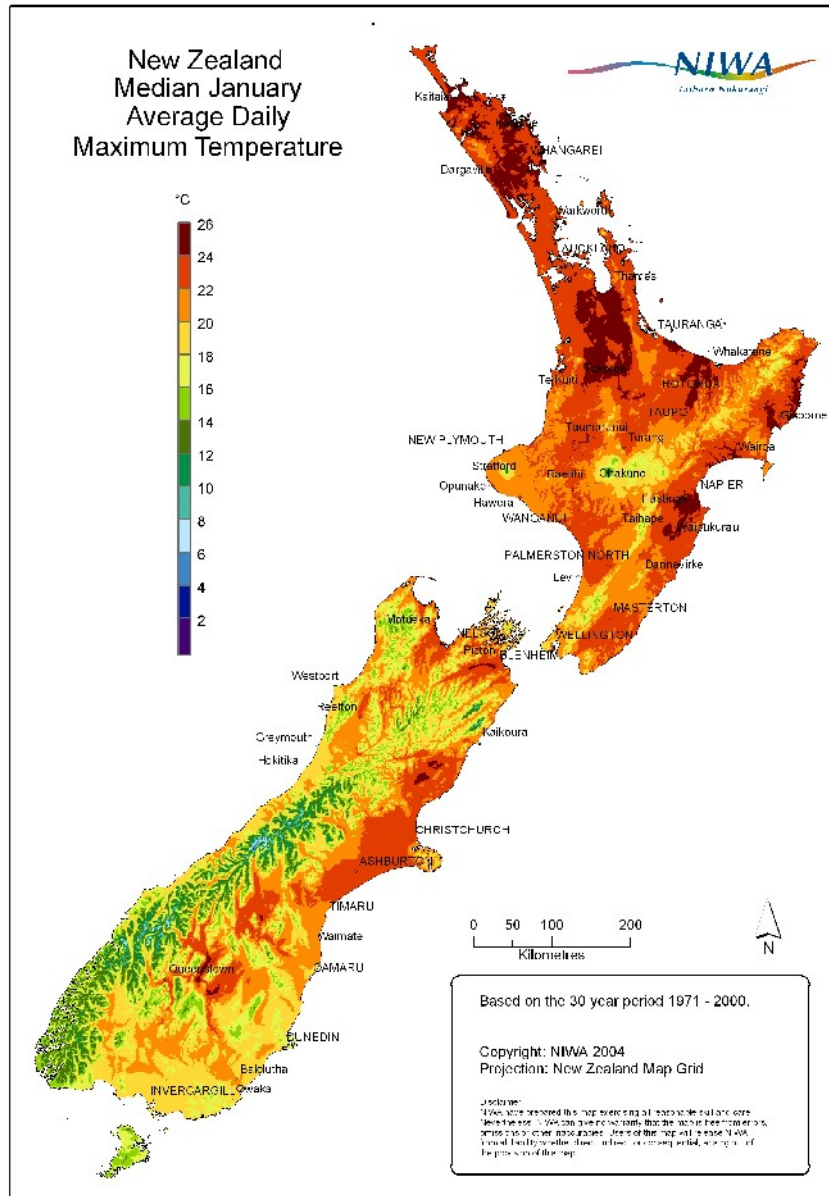
Oxidation of N₂O (Reaction (5)) is the major source of stratospheric NO_x (NO+NO₂+NO₃+HNO₃+2N₂O₃+HNO₄+ClONO₂+BrONO₂) and hence NO_x (Minschwaner et al., 1993).



N₂O concentrations are predicted to continue to increase over the coming century due to anthropogenic surface emissions, mostly attributed to agricultural nitrogen fixation (IPCC, 2001; WMO, 1999). Trends in the concentration of atmospheric N₂O for the period 1980 to 1988 have been estimated to be +0.25±0.05% per year (IPCC (2001), p253). Zander et al. (1994) quote a trend in N₂O of +0.33±0.04 %

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 (h.struthers@niwa.co.nz)





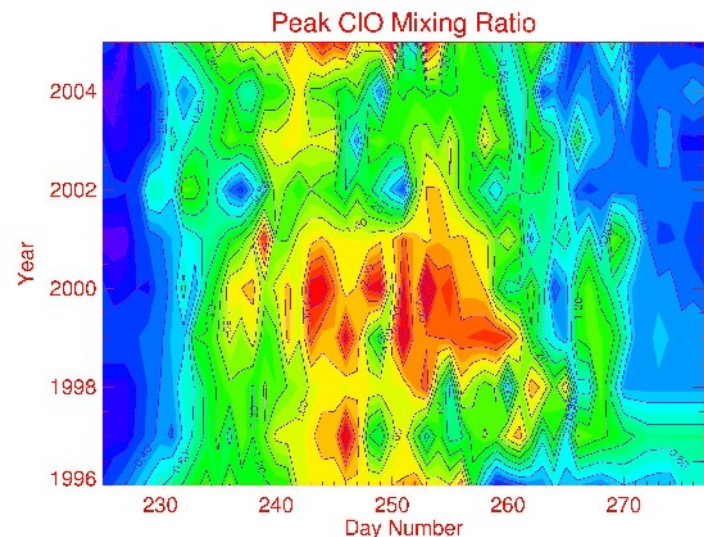
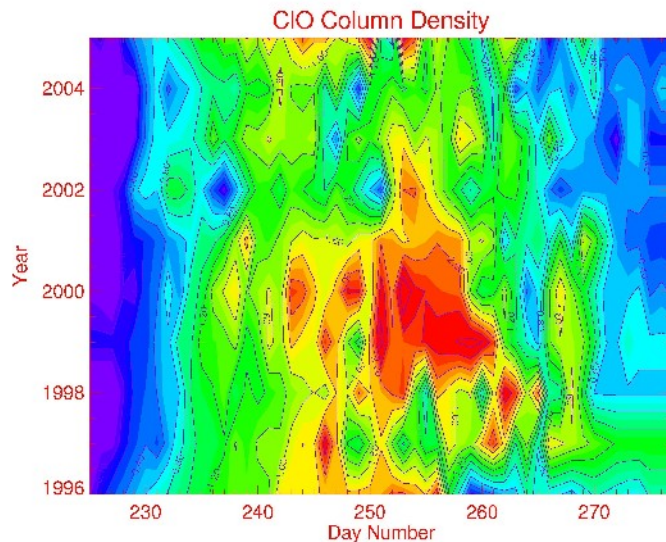
Lauder Measurements at the Antarctic NDACC Station, Dec 2007

Instrument & Period	Parameter	Cooperating Institutions	Comments
Dobson Ozone 1987 -	Column ozone	NOAA/GMD, Boulder, USA	Direct sun obs. Winter moon obs. (when visible)
IR Interferometer (Bomem FTIR) 1992 - 96 (Bruker FTIR) 1997 -	Column HCl, HNO ₃ , ClONO ₂ , HF, CFCs, CO & GHGs, etc	University of Denver, Denver, USA	Mid-IR (windows in 2-12μ region). SFIT2 profiles
UV-Vis DOAS spectrometers 1982 -	Column NO ₂ Column OClO Column BrO Column ozone Off-Axis BrO	University of Heidelberg	340-490 nm Halogens start 1993, Off-axis 1998.
ClO microwave radiometer 1996 -	ClO	SUNY, Stony Brook, New York, USA	278 GHz and 1 Day integration time

CIO Time Series, 1996-2005. Measured with the Stonybrook Microwave Radiometer at Scott Base

Contour plots of column density and peak mixing ratio vs. year and day of year

The Figures below show CIO column density and peak mixing ratio as a function of time, with day of year on the x-axis and year on the y-axis. The maximum values of CIO occur during the period of days 240-260, and the onset of CIO is less variable than its annual decline. The multi-year record shows that the highest values of CIO occurred in 1999-2000, that values were consistently low in 2002-2004, but rebounded significantly in 2005. In addition, there is a hint of the annual decline occurring earlier over the course of the 10 years.



The peak mixing ratio plot shows essentially the same features as the column density.

